

## 1 Overview

The African superplume, is widely considered to have caused the ~30Ma volcanism at the Ethiopian traps. The contribution of more localised upwellings to African volcanism, that may or may not originate below the mantle transition zone is the subject of considerable debate.

Thermochemical conditions impart control on the depth at which mantle materials undergo phase changes, causing impedance contrasts. Observations of seismic discontinuities from the mantle transition zone (MTZ) and below can therefore provide insight into the variable thermochemical nature of upwellings beneath Africa.

Here we present observations of seismic discontinuities beneath Africa obtained from a compilation of P-to-s (Pds, PPs, and PKPs) receiver functions derived from publicly available seismograph networks across Africa from 1990-2019.

We capitalise on a new high-resolution P-wave absolute velocity model for the African continent (Boyce et al., in prep.) to migrate our receiver functions to depth prior to common conversion point (CCP) stacking.

We interrogate our receiver function CCP stacks for MTZ discontinuity topography at a range of frequencies and for discontinuities at mid-mantle depths (~1000km) to understand better the variable thermochemical nature of mantle upwellings that have contributed to recent volcanism across Africa.

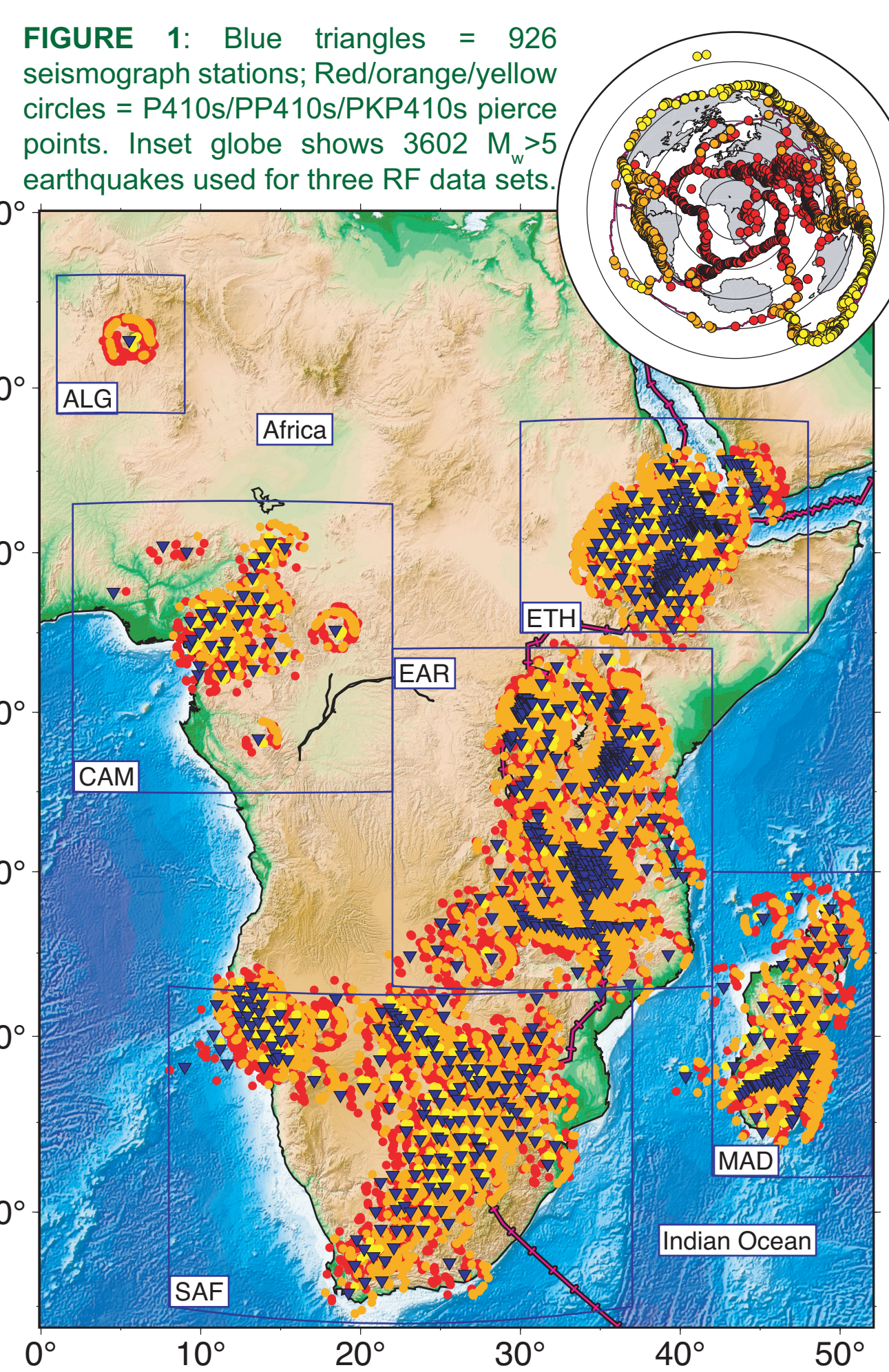
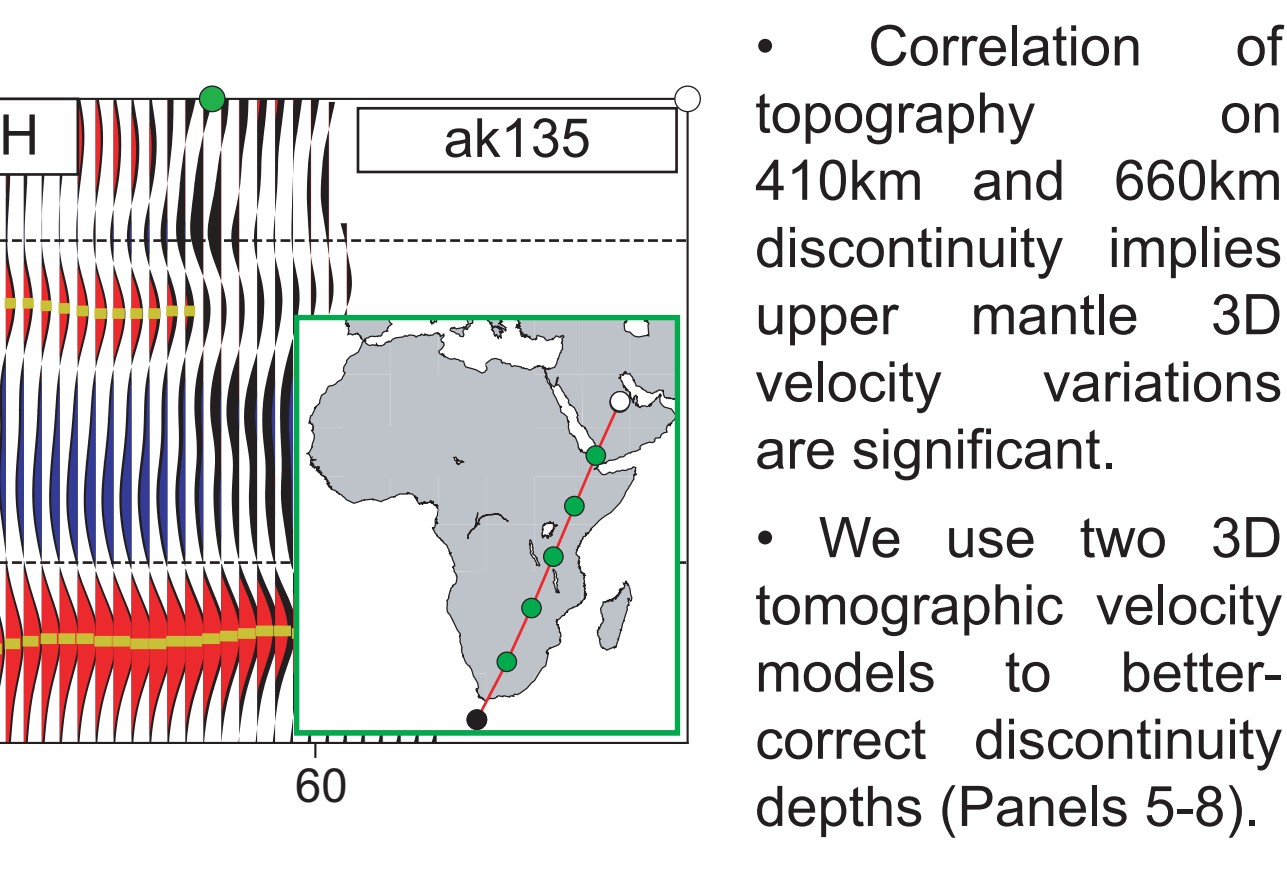
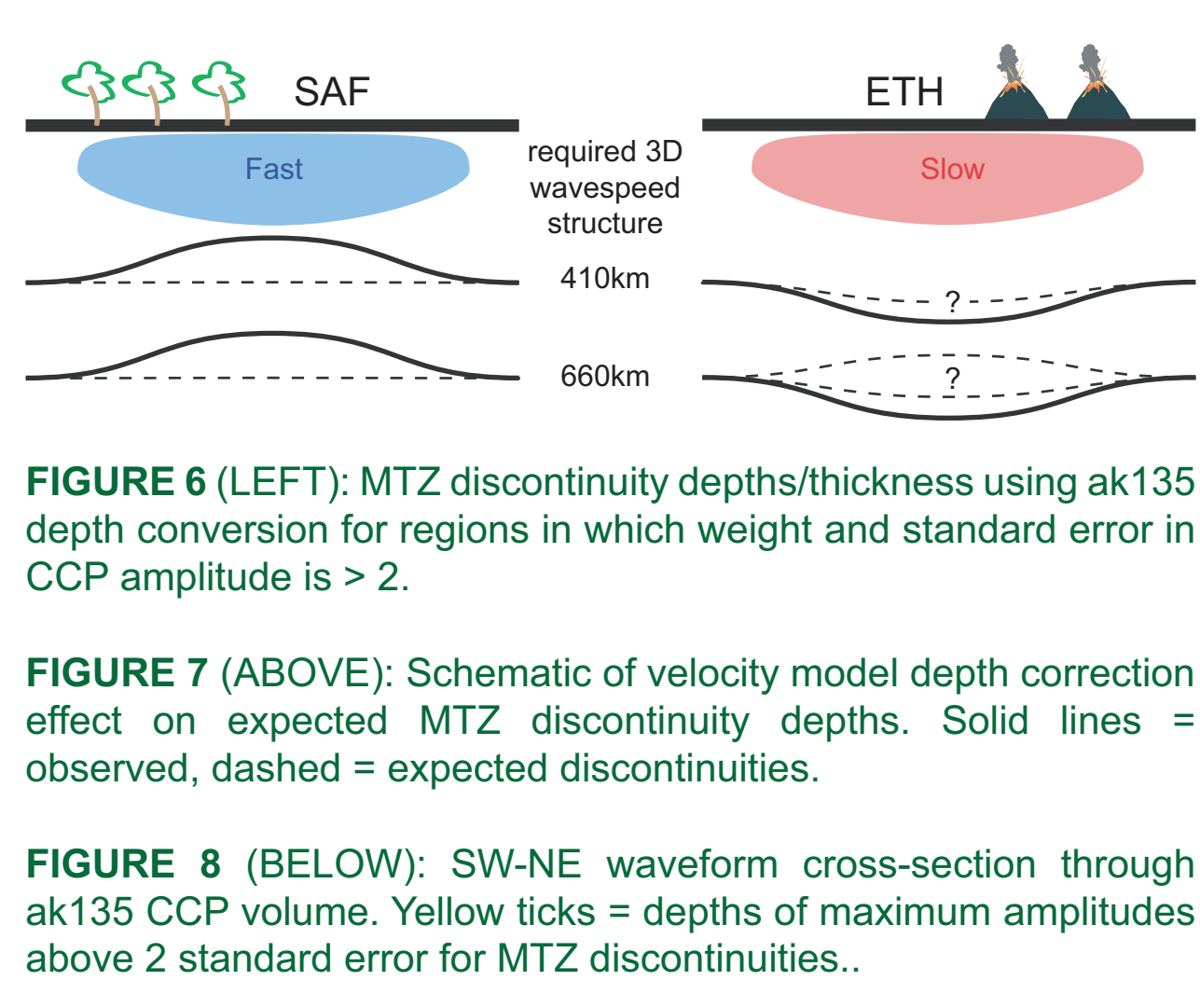
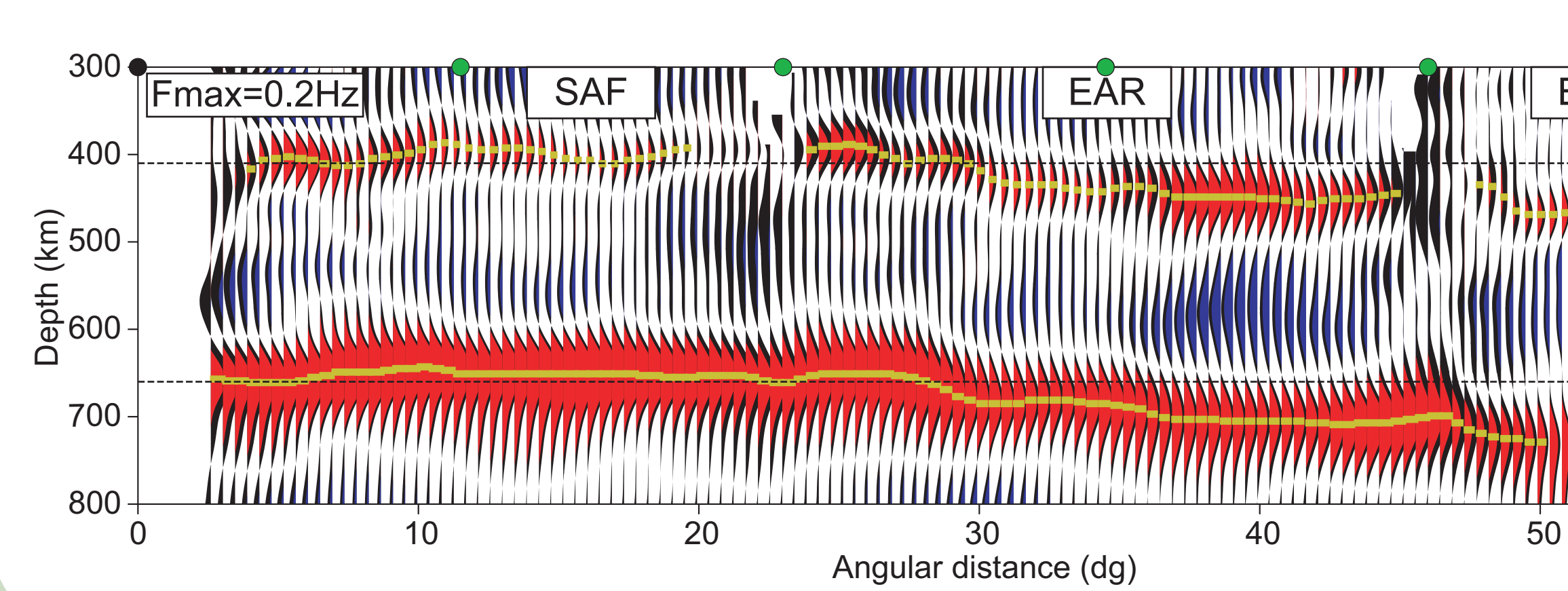
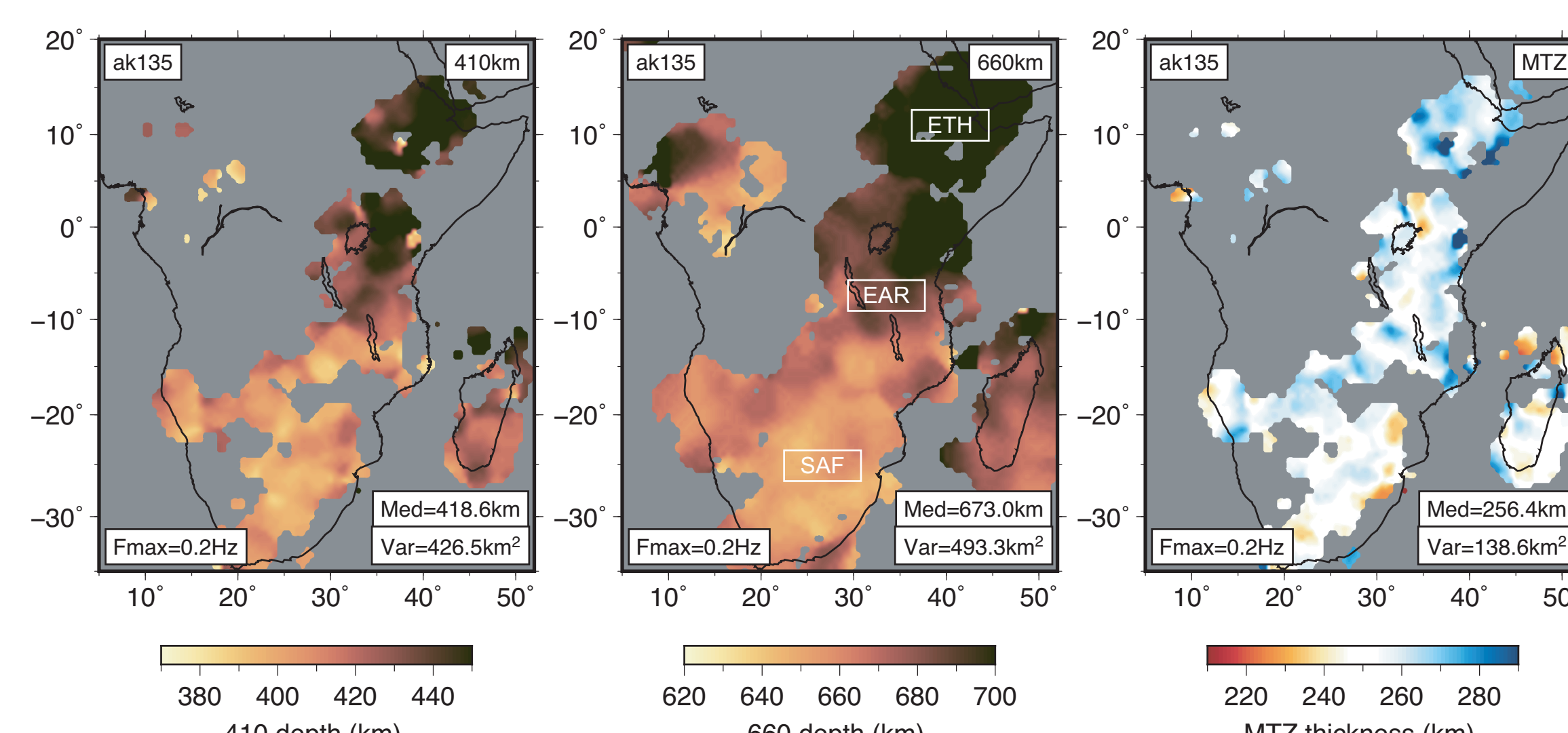


FIGURE 1: Blue triangles = 926 seismograph stations. Red/orange/yellow circles = P410s/PP410s/PP600s/PP1000s. Inset globe shows 3602 M<sub>w</sub>>5 earthquakes used for three RF data sets.

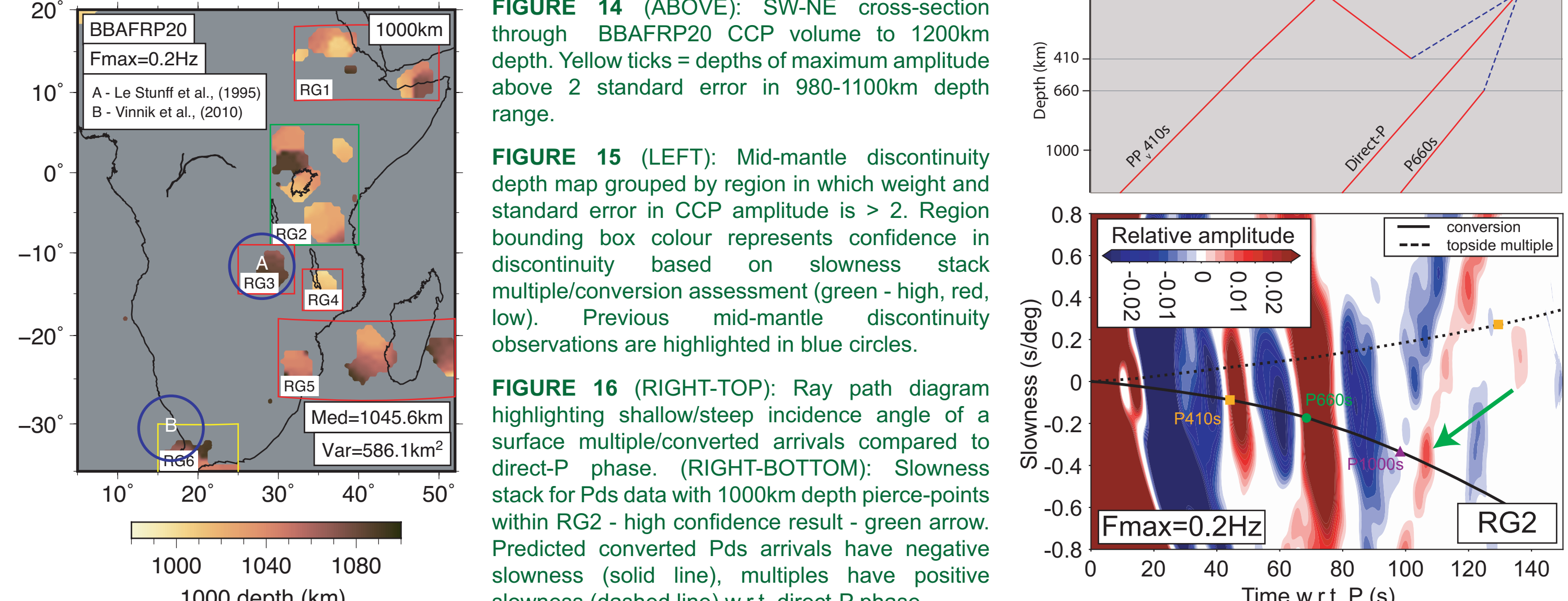
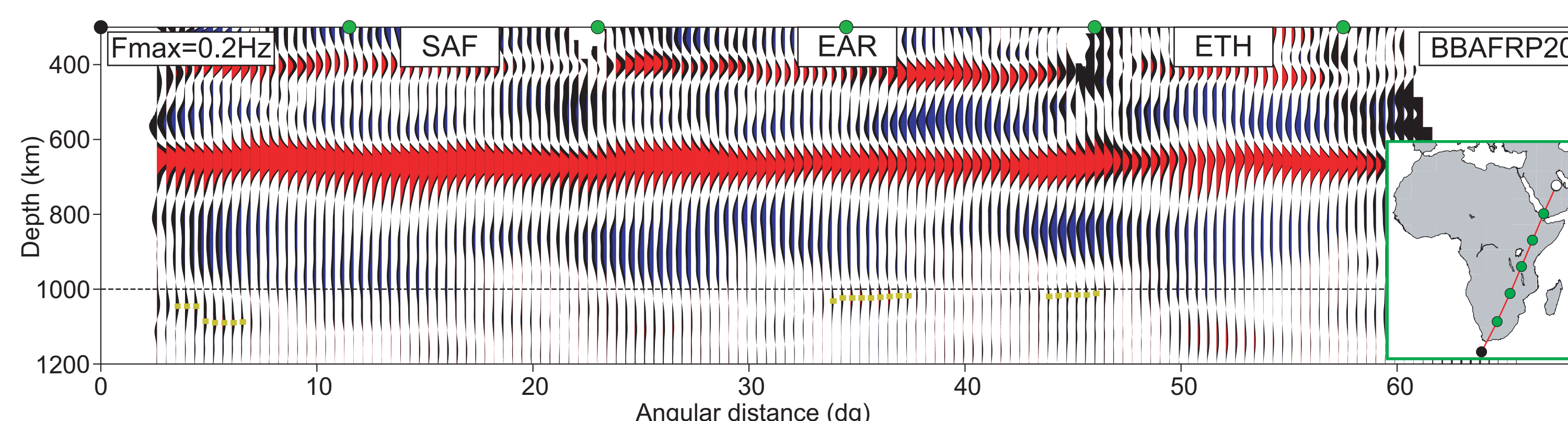
## 4 CCP stack 1D velocity model - ak135



Correlation of topography on 410km and 660km discontinuity implies upper mantle 3D velocity variations are significant.

We use two 3D tomographic velocity models to better-correct discontinuity depths (Panels 5-8).

## 7 Mid-mantle observations



Observations of Pds conversions beneath EAR (Figure 14 and RG2 - Figure 15) at ~1025km depth.

Slowness assessment (Figure 16) highlights coherent arrival for observations in RG2 (green arrow), lower confidence in mid-mantle conversions elsewhere.

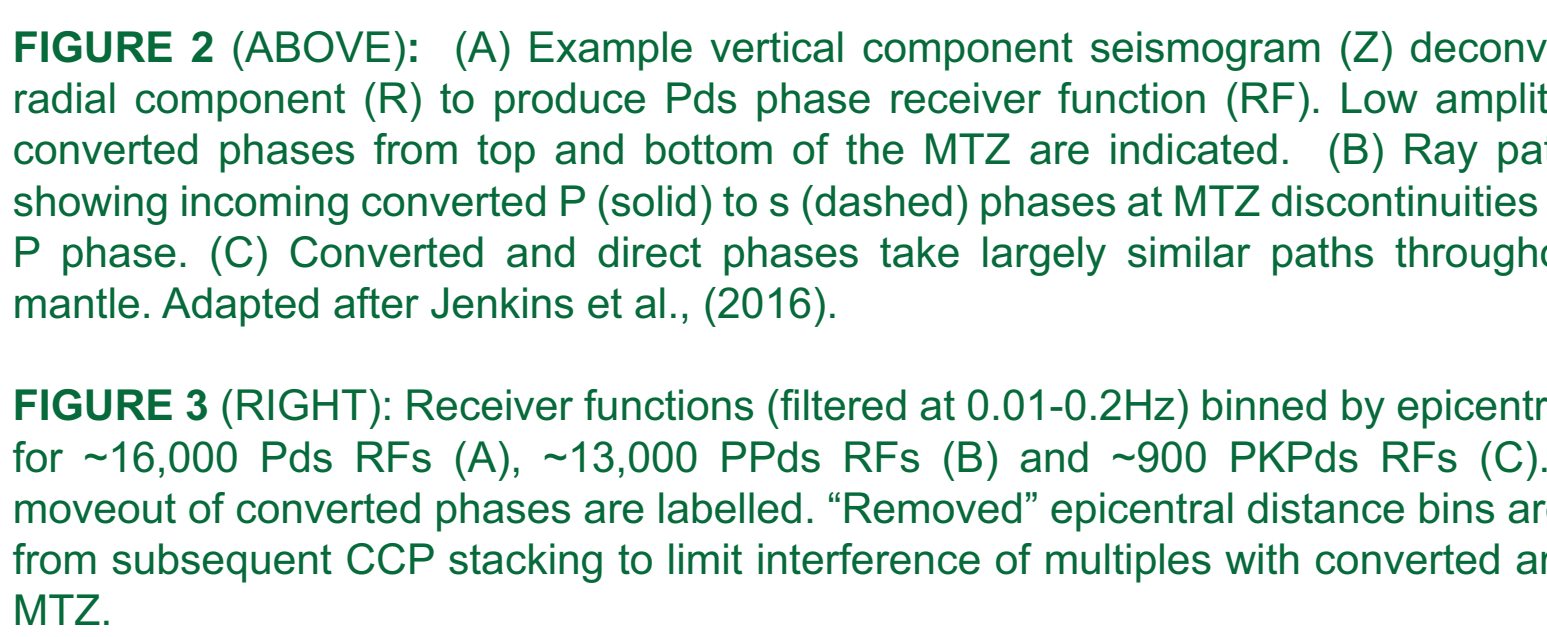
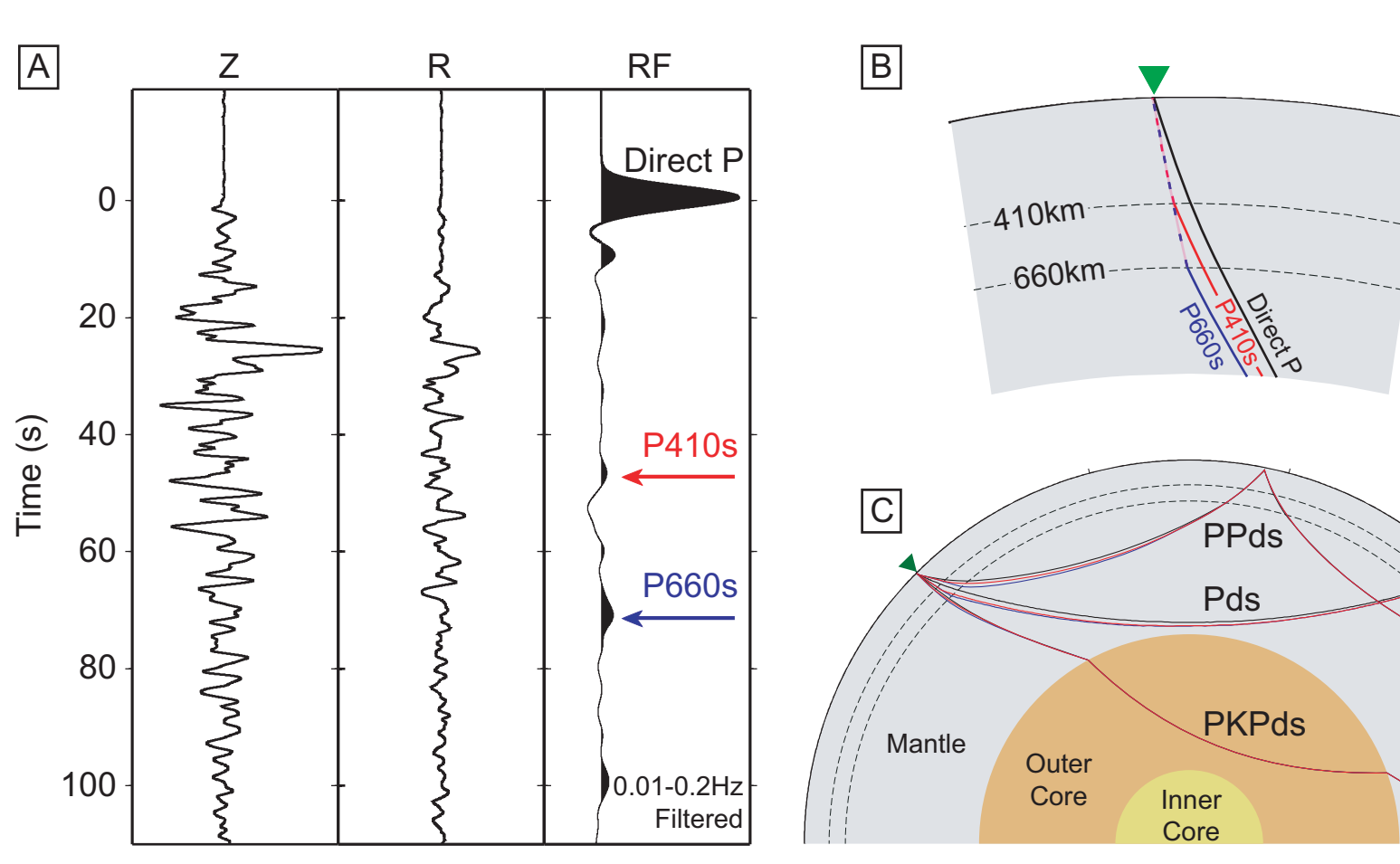
Converted arrival from ~1025km depth in RG2 collocated with slow wavepicks in BBAFRP20 at 1000km depth - green arrow (Figure 17).

A warm MTZ displaying depressed 410km and 660km discontinuities, underlain by a ~100km depth discontinuity may indicate entrained heterogeneity in a plume.

Chemically distinct plume material may derive from recycled basaltic material or a deep sourced LLSVP reservoir (Jenkins et al., 2017).

FIGURE 17 (ABOVE): P-wave absolute velocity structure for BBAFRP20 (Boyce et al., in prep) at 1000km depth, plotted as deviation (dVp %) from ak135.

## 2 Receiver function calculation

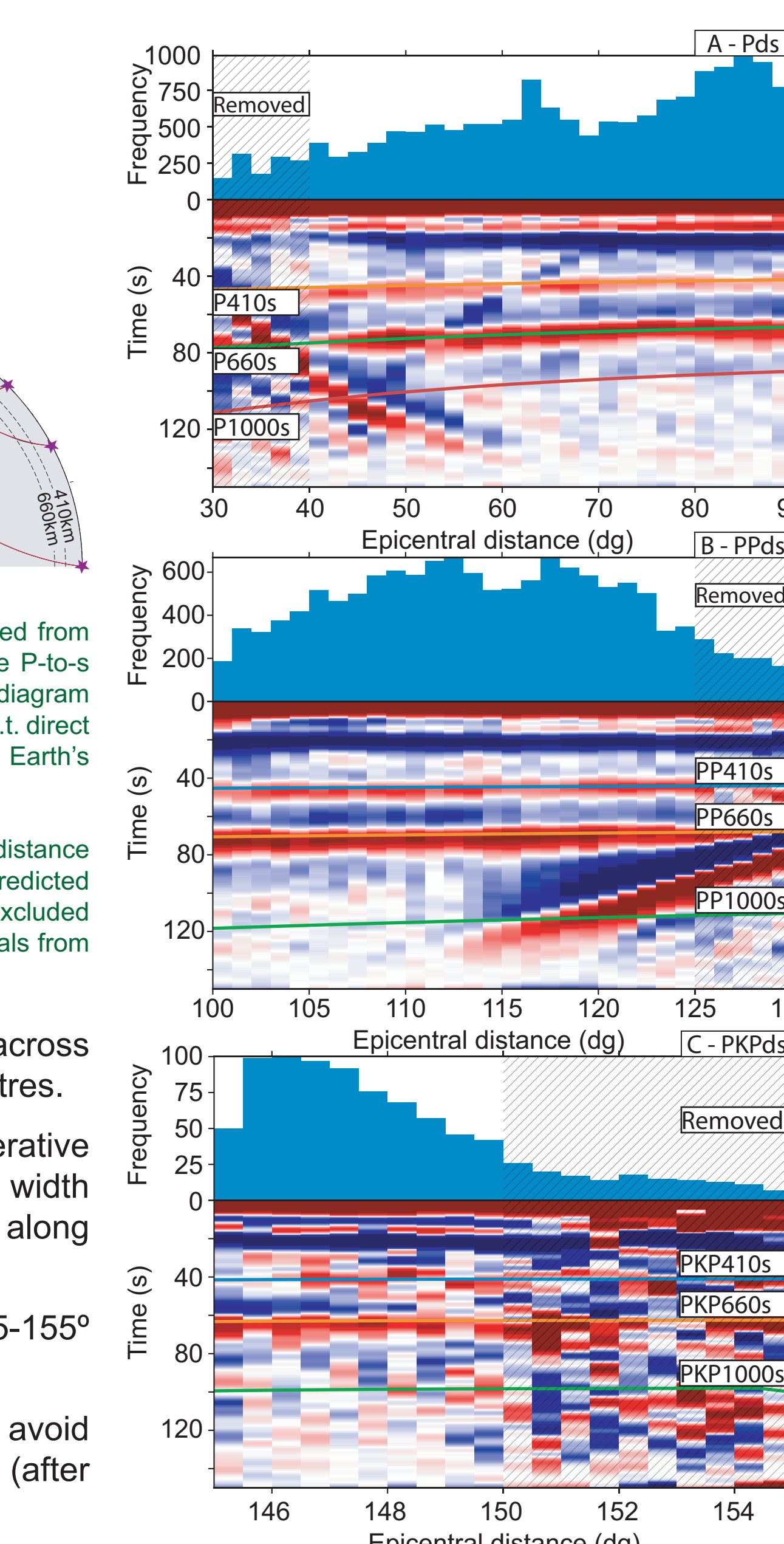


Waveform data set sourced from publicly available stations across Africa (1990-2019) through IRIS, GEOFON and RESIF data centres.

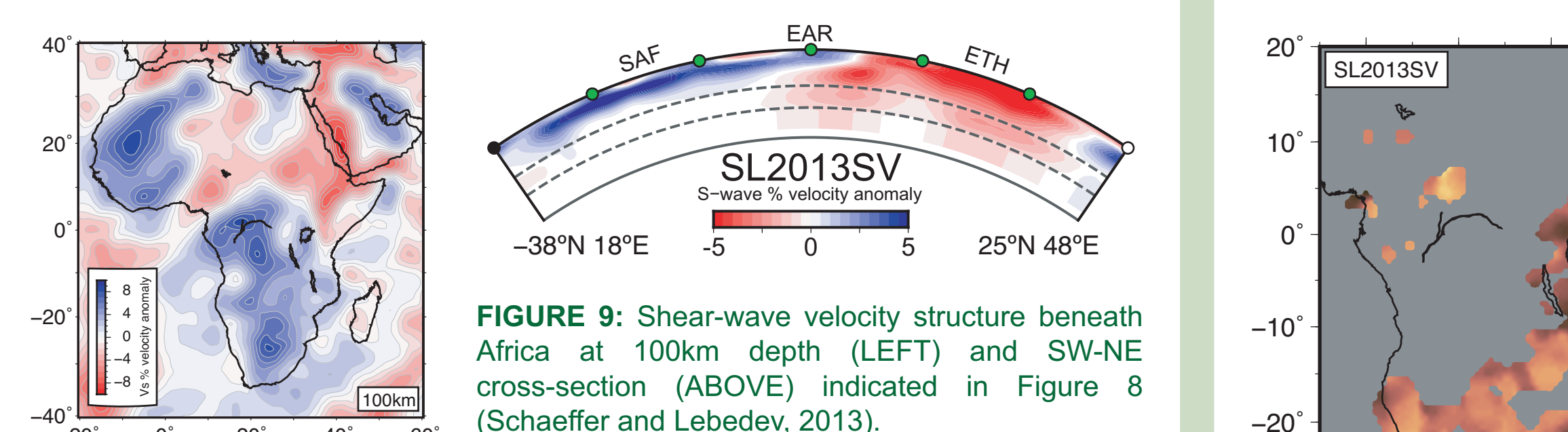
Receiver functions (RFs) are calculated using time-domain iterative deconvolution (Ligorria and Ammon, 1999) initially built with 5s width Gaussian pulses to obtain a representation of earth structure along incoming ray paths (Figure 2).

Pds, PPs, PKPs phases recorded at 30-90°, 100-130°, 145-155° epicentral distance respectively (Figure 3A-C).

We impose epicentral distance restrictions upon these data to avoid phase interference (Figure 3) and apply strict automated QC (after Cottaar and Deuss, 2016).



## 5 3D mantle velocity structure



Continental CCP stacks computed using depth corrections from SL2013SV and a new V<sub>p</sub> model BBAFRP20 (Figures 9, 10).

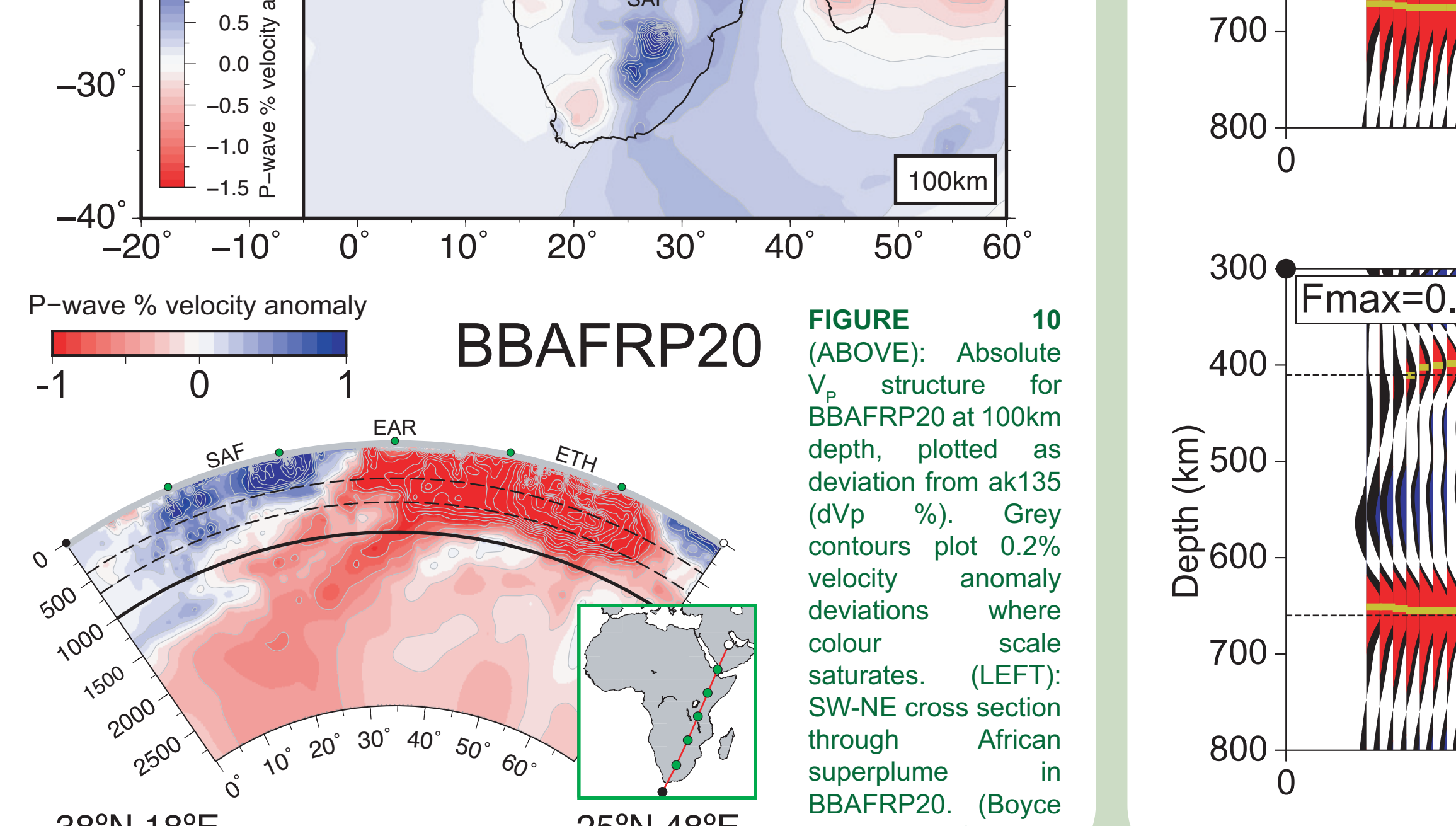
Global P-phase pick database (1964-2007), inverted for absolute V<sub>p</sub> structure using an adaptively parameterized, LSQR inversion scheme (Li et al., 2008).

Supplemented with >87,000 absolute P-wave delay times from African temporary networks (1990-2019) calculated using AARM (Boyce et al., 2017).

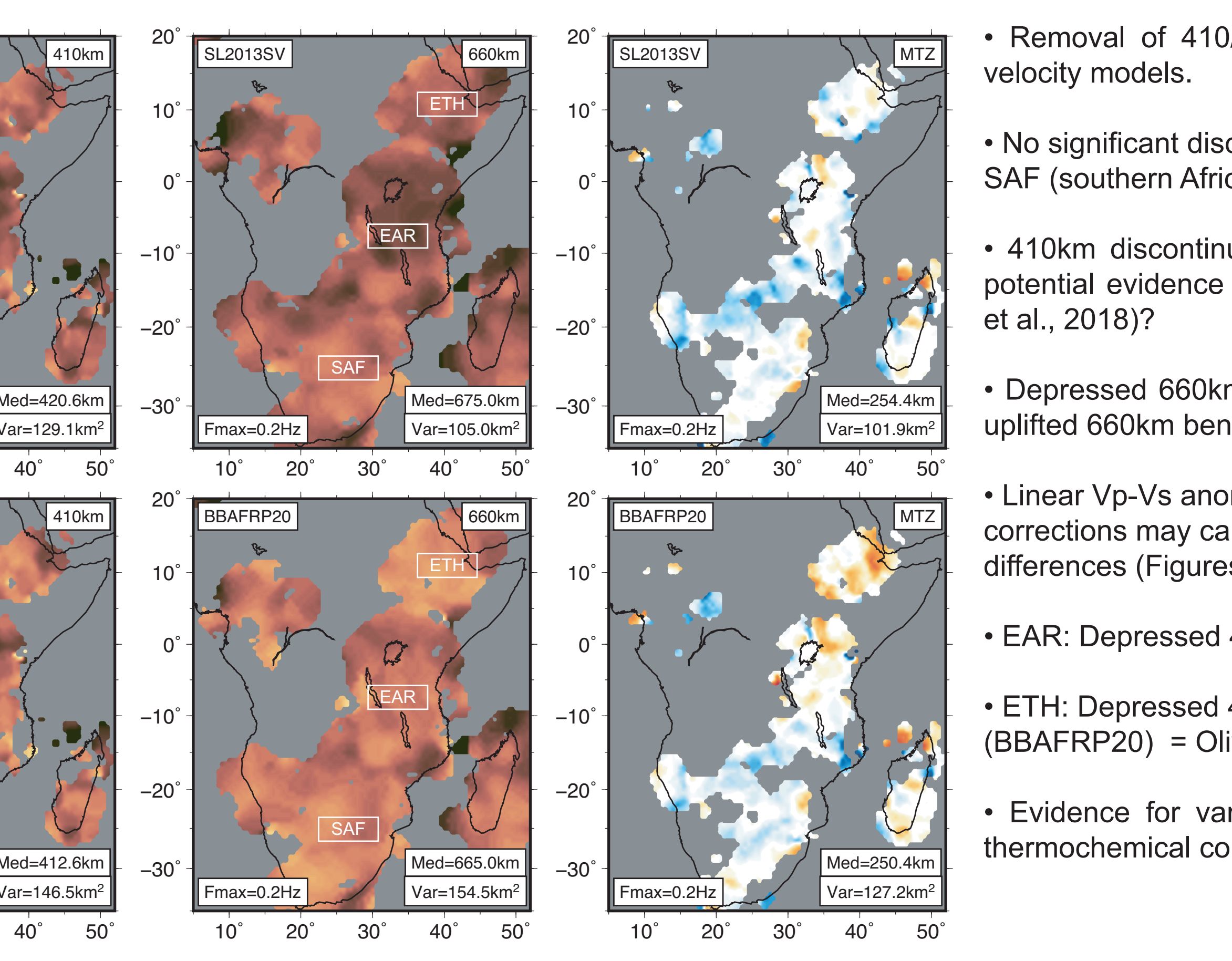
Upper mantle V<sub>p</sub> structure of southern and eastern Africa well resolved beneath available stations (Figure 1).

Slow V<sub>p</sub> anomaly extends from CMB below southern Africa across MTZ below EAR. Evidence for additional sub-vertical slow V<sub>p</sub> anomaly below Ethiopia/Somalia.

Velocity anomalies are anticorrelated between V<sub>s</sub> and V<sub>p</sub> (SL2013SV and BBAFRP20) in upper mantle below EAR.



## 6 CCP stack 3D velocity models - SL2013SV & BBAFRP20



Removal of 410/660km correlation using 3D velocity models.

No significant discontinuity topography beneath SAF (southern Africa).

410km discontinuity not visible in Cameroon, potential evidence for water filled MTZ (Buchen et al., 2018)?

Depressed 660km beneath EAR, evidence for uplifted 660km beneath ETH (BBAFRP20).

Linear Vp-Vs anomaly scaling in depth corrections may cause EAR MTZ topography differences (Figures 9, 10)?

EAR: Depressed 410&660km = Garnet?

ETH: Depressed 410km, uplifted 660km (BBAFRP20) = Olivine?

Evidence for variable response of 660km to thermochemical conditions beneath EAR-ETH.

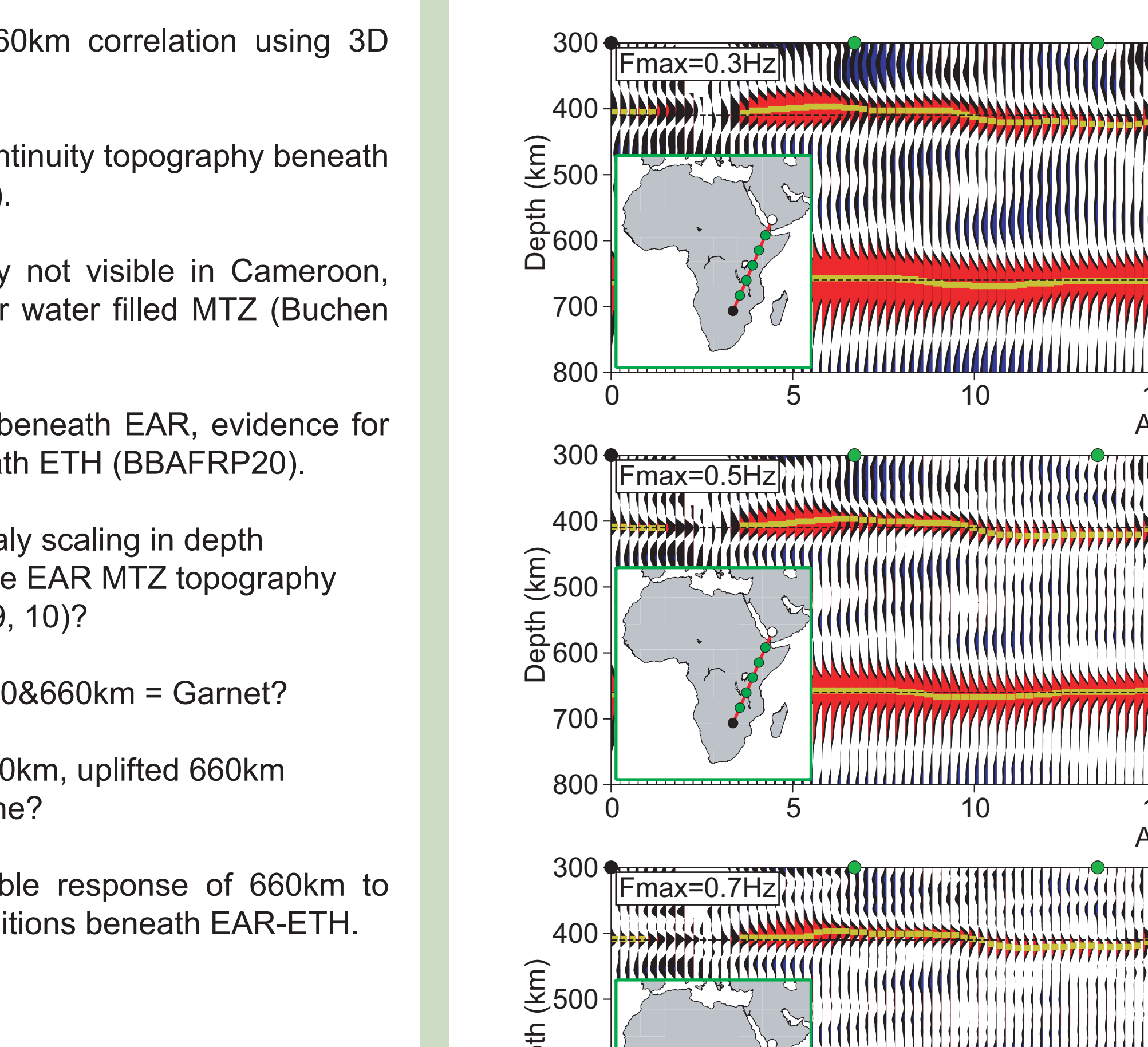
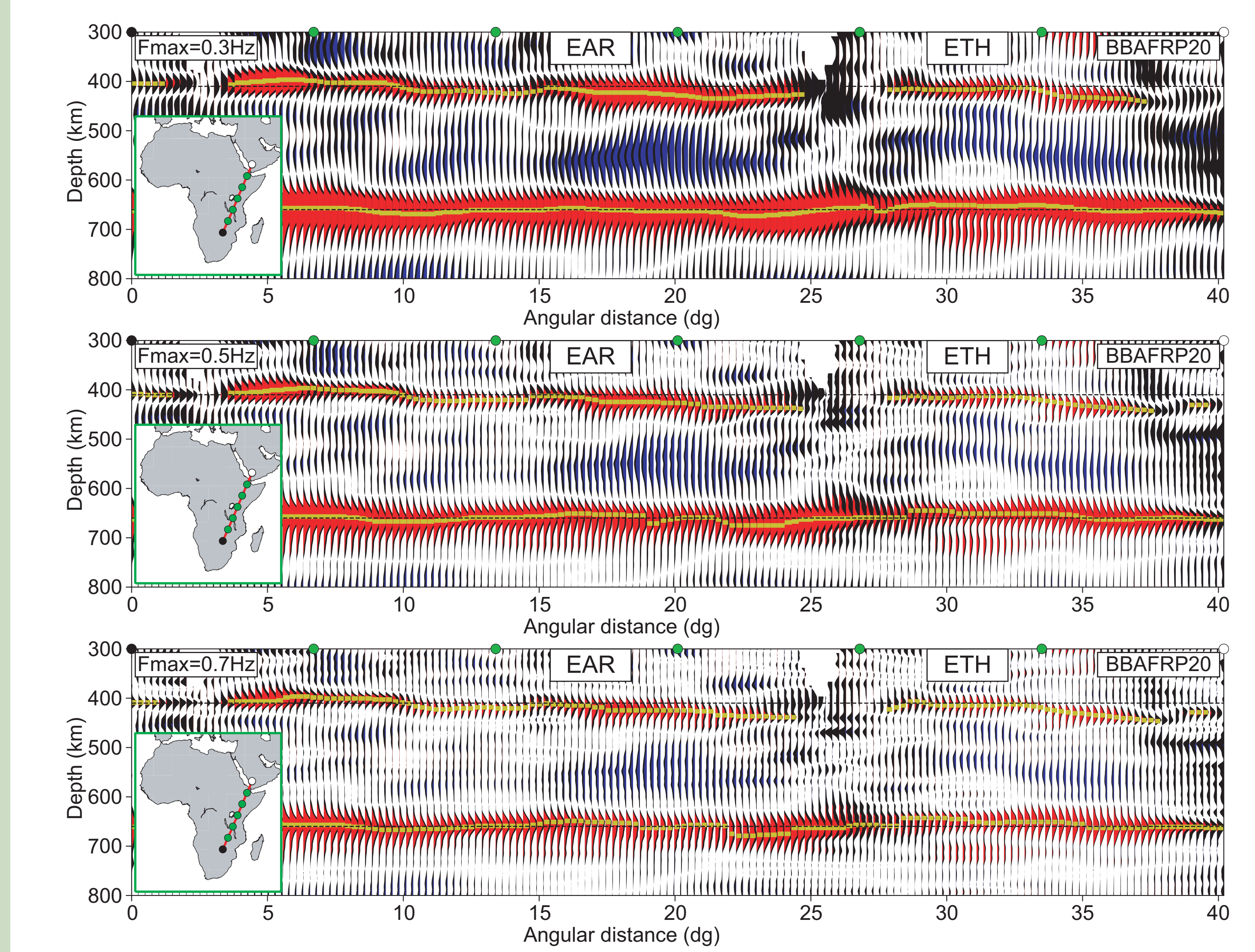


FIGURE 11 (ABOVE): MTZ discontinuity depths/thickness using SL2013SV and BBAFRP20 depth conversion for regions in which weight and standard error in CCP amplitude is > 2.

FIGURE 12 (RIGHT): Interpretation guide for MTZ discontinuity depths for a warm upper mantle. Opposing clapeyron slopes of ringwoodite and majorite phase transitions to perovskite can produce 660km discontinuity topography.

FIGURE 13 (BELOW): SW-NE waveform cross-section through SL2013SV (TOP) and BBAFRP20 (BOTTOM) CCP volumes. Yellow ticks = depths of maximum amplitudes above 2 standard error for MTZ discontinuities.

## 8 Frequency analysis



660km discontinuity depth is uplifted across three frequency bands beneath ETH.

660km discont. appears shallower at high frequencies below EAR, evidence for two peaks/mineral phase transitions between ~650-700km depth.

Presence of both ringwoodite and majorite phase transitions to perovskite below EAR?

## 3 Depth conversion and CCP stacking

Delay times of S-phases can be converted to a discontinuity depth using an upper mantle velocity model.

To account for the spread of conversion points (Figure 1) and the increasing Fresnel zone width with depth (Figure 4), we use weighted common conversion point (CCP) stacking (e.g., Cottaar and Deuss, 2016).

RF amplitudes are back propagated along ray paths and stacked using a radially decreasing weighting at grid points within 2 Fresnel zone half widths (Figure 4).

Initially we compute discontinuity depths using ak135 (Kennett et al., 1995 - Panel 4).

Subsequently we use an upper mantle shear wave model - SL2013SV (Schaeffer and Lebedev, 2013) and a new V<sub>p</sub> model for Africa - BBAFRP20 (Boyce et al., in prep. - Panel 5).

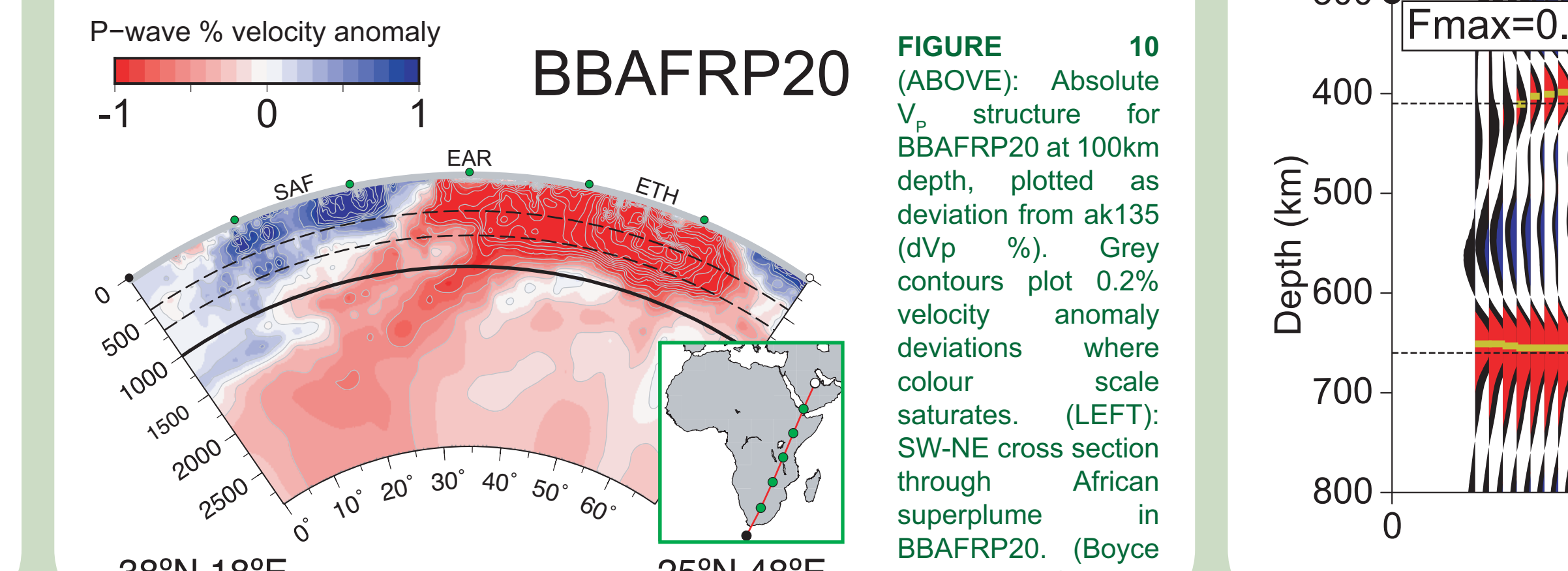
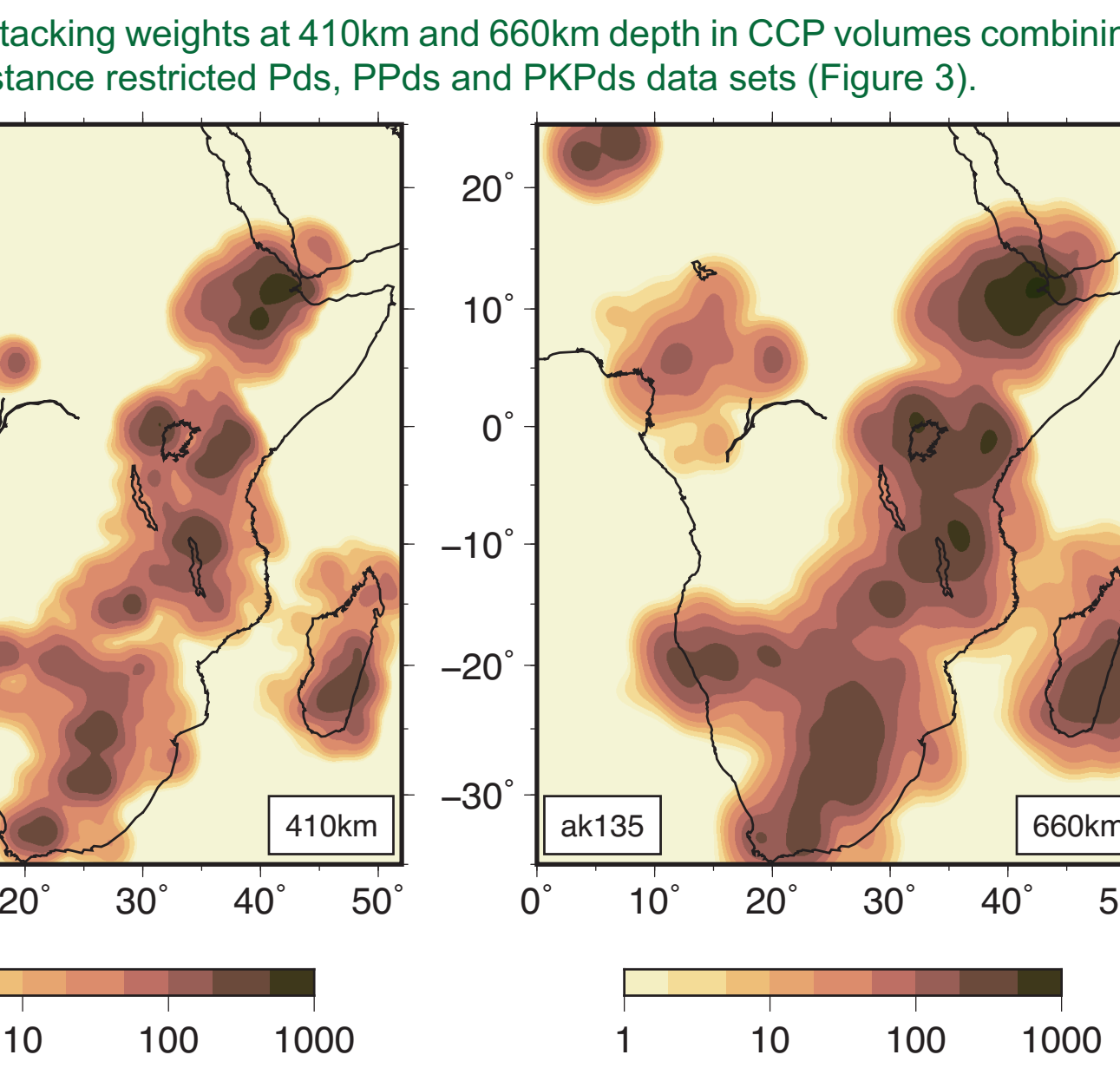


FIGURE 10 (ABOVE): Absolute Vp structure for BBAFRP20 at 100km depth, plotted as deviation from ak135 (dVp %). Grey contours plot 0.2% velocity anomaly deviations where colour scale saturates. (LEFT): SW-NE cross section through African superplume in BBAFRP20. (RIGHT): Absolute Vp structure for BBAFRP20 at 100km depth, plotted as deviation from ak135 (dVp %).

## 9 Key points

BBAFRP20 P-wave tomographic model provides appropriate time-to-depth correction for RF stacks across Africa.

Coherent 410km not visible in Cameroon, water in the MTZ?

Variable character of warm MTZ beneath east Africa; a variable source in depth or deep mantle reservoir(?)

EAR: Frequency dependent, Garnet(?) dominant MTZ, 1000km discont.

ETH: Olivine(?) dominant MTZ, no 1000km discont.

Support for variable source of volcanism in east Africa - two plumes?

**Hypothesis:** The African superplume samples chemically distinct material from LLSVP whilst a vertical plume beneath Ethiopia/Afar does not?

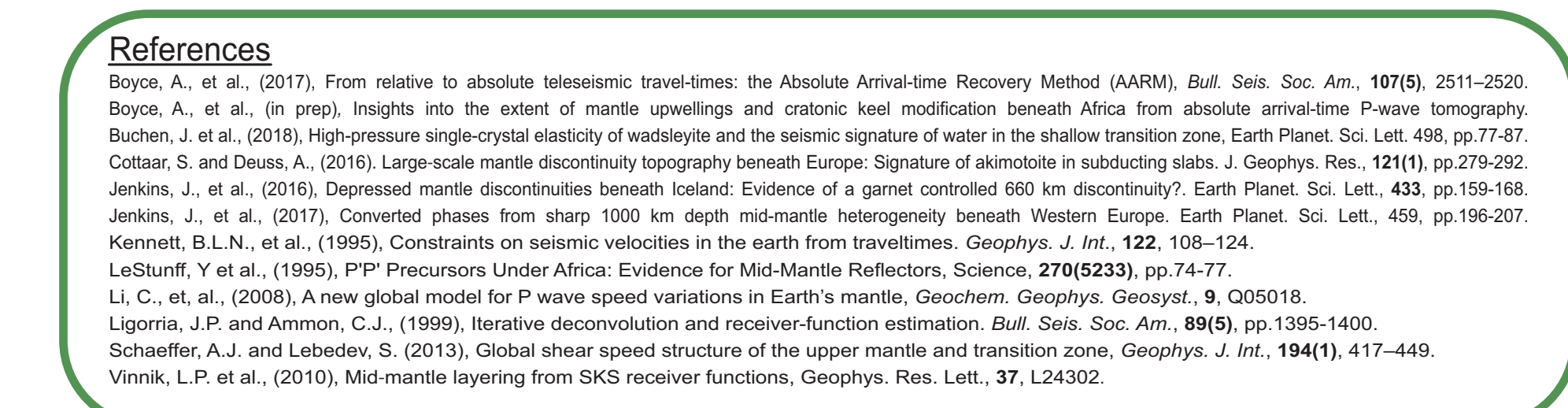


FIGURE 19: Schematic interpretation of variable character of mantle upwellings beneath east Africa.